



## Thickness of marine Holocene sediment in the Gulf of Trieste (Northern Adriatic Sea)

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**Abstract.** We use various geophysical datasets (multibeam and singlebeam echosounder data, sub-bottom profiling Chirp and sonar data and very high resolution boomer seismic data) along with published  
20 sedimentological data and depth data from nautical charts in order to create models of the depth of the seafloor and the base of Holocene marine sediment in the Gulf of Trieste. The two models are later used in order to calculate the thickness of marine Holocene sediment which has been depositing on the Late Pleistocene alluvial plain since the Holocene transgression in the Italian and Slovenian part of the gulf. Thicker Holocene marine sedimentary sequences averaging at around 5 meters are characteristic for the SE part of the gulf. In other parts  
25 of the gulf the Holocene marine sedimentary cover is very thin or even absent, except in close proximity of the shoreline and fluvial sediment sources, in the area of the Trezza Grande paleodelta and above topographic depressions of the Late Pleistocene base. The presented datasets available from the OGS SNAP data repository (<https://doi.org/10.6092/6ad9b1e6-c977-cec9-8a2d-db10c7f90adc>) represent a valuable reference for a wide variety of research disciplines dealing with the dynamic Earth system in the Gulf of Trieste and could use as a  
30 valuable tool for designing sampling and geophysical campaigns in the studied area.

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### 1 Introduction

40 The Gulf of Trieste is a shallow gulf with an average depth at around 20 meters located in the northernmost part of the Adriatic Sea where it is surrounded by Italian, Slovenian and Croatian coasts (Fig. 1). After the Last Glacial Maximum with the onset of the Holocene transgression, marine sediment started depositing approximately 10,000 years ago in this area (Marocco, 1991; Ogorelec et al., 1991, 1997; Lambeck et al., 2004; Covelli et al., 2006; Ogrinc et al., 2007; Trincardi et al., 2011b).



Early research of Holocene marine sediment in the Gulf of Trieste was mainly limited to sedimentological and geochemical investigations of data acquired with a relatively small number of cores and wells located onshore and offshore the gulf in the 80's and 90's (Ogorelec et al., 1981, 1991, 1997; Marocco et al., 1984; Marocco, 1989, 1991; Faganeli et al., 1991; Gordini et al., 2002; Covelli et al., 2006; Ogrinc et al., 2007) and investigations for the Italian Marine Geological Map 1:250.000 of Venice (Trincardi et al., 2011a, 2011b). The development of cost-effective geophysical methods in recent years resulted in a number of geophysical surveys undertaken in the last decade that were focused on high-resolution geophysical investigation of the seafloor and the sub-seafloor geological structure of the gulf (Gordini et al., 2003, 2004, 2006, Gordini, 2007, 2009; Romeo, 2009; Trincardi et al., 2011a, 2011b; Kolega and Poklar, 2012; Slavec, 2012; Zampa et al., 2015; Zecchin et al., 2015; Trobec, 2015; Trobec et al., 2016, 2017). These extensive datasets along with additional unpublished data spanning over a mayor part of the gulf allowed us to assess the thickness of the Holocene marine sediment in the Gulf of Trieste.

This work aims to present the first comprehensive model of the distribution and thickness of Holocene marine sediment in the Slovenian and Italian parts of the Gulf of Trieste. The model is derived from geophysical, core/well, nautical chart data and in parts complemented with the published Holocene marine thicknesses from Trincardi et al. (2011b).

## 1.1 Geological setting of the study area

A simplified geological overview of the surroundings of the Gulf of Trieste consists of Quaternary alluvial sediment of the Friuli plain on the north, Cretaceous – Paleogene carbonates of the Classical Karst on the northeast and Paleogene marls and sandstones of the flysch of the Istria peninsula and Trieste coastline on the southern and eastern part of the gulf (Placer, 2015; Biolchi et al., 2016; Jurkovšek et al., 2016). A similar geological sequence can be observed in geophysical data offshore the gulf area where the carbonate platform is followed by a flysch succession which is overlain by a few hundred meters of Quaternary sediment deposited during the transgressive-regressive cycles (Busetti et al., 2010a, 2010b; Cimolino et al., 2010; Vrabec et al., 2014). The youngest sedimentary sequence in the Gulf of Trieste is represented by Holocene marine sediment which has been depositing for the last 10,000 years since the onset of the Holocene transgression following the Last Glacial Maximum (Ogorelec et al., 1981, 1997; Lambeck et al., 2004; Covelli et al., 2006; Trincardi et al., 2011b; Zecchin et al., 2015). Many authors suggest that the base on which marine sediment has been depositing, represents relict continental-paralic sedimentary environments predating the Holocene sea transgression in the Gulf of Trieste (Ogorelec et al., 1981, 1991, 1997; Marocco, 1989; Lambeck et al., 2004; Covelli et al., 2006; Trincardi et al., 2011b; Slavec, 2012; Zecchin et al., 2015; Trobec, 2015; Trobec et al., 2017). The sediments of the Holocene marine transgression provided a thin and discontinuous deposit occurring on an erosive surface of the Late Pleistocene sediments due to subaerial exposure (Trincardi et al., 2011a).

## 2 Data used and modelling of the different surfaces



Numerous geophysical surveys were conducted between 2000 and 2015 in the Gulf of Trieste by the Istituto Nazionale di Oceanografia e di Geofisica Sperimentale (OGS) and by Harpha Sea d.o.o in cooperation with the Department of Geology of the University of Ljubljana (Fig. 4 and 5, Tables 2 and 3).

Geophysical data was correlated and calibrated with published core/well data (Table 1 and Figs. 2 and 3 (Ogorelec et al., 1981, 1991, 1997; Marocco et al., 1984; Marocco, 1989; Gordini et al., 2002; Covelli et al., 2006; Romeo, 2009; Trincardi et al., 2011b; Zecchin et al., 2015). The used cores/wells usually document a transition from an alluvial to paralic and later marine sedimentary environment. However, in some cores the marine or paralic sedimentary sequence is absent or marine and paralic sediments are deposited directly on the bedrock (Figs. 2, 3 and 8(d)).

Geophysical datasets along with previously published core/well data allowed us to create a model of the distribution and thickness of Holocene marine sediment in the Italian and Slovenian parts of the gulf. The different datasets and methods used to create each model are described in detail in the subsequent subchapters.

## 2.1 The bathymetric model

In order to create the bathymetric model of the Gulf of Trieste we used multibeam and singlebeam sonar data complemented with high resolution single channel seismic (boomer and chirp) (Fig. 4, Tables 2 and 3). In addition, water depths of 28 points in the Gulf of Panzano were determined from the Slovenian nautical chart of the Gulf of Trieste (Ministry of Transport of the Republic of Slovenia, 2005) and the “Da Punta Tagliamento a Pula” nautical chart (Istituto Idrografico della Marina, 2004). We also used depths of published cores/wells in the Gulf (Table 1, Fig. 2). The used datasets are depicted in Fig. 4. In order to constrain the model, we used the coastline contour for which we assumed a seafloor depth value equal to 0 m above sea-level.

Multibeam sonar data included data acquired during the reflection seismic surveys in the central and Slovenian part of the gulf in the years 2009 and 2013 (Table 2). In addition, a multibeam based bathymetric model of Slovenian territorial waters (Slavec, 2012) acquired by Harpha Sea d.o.o. was also used. Smaller areas surveyed by OGS by multibeam sonar also include three areas in front of the Grado-Marano lagoon: The Porto Buso inlet, Morgo inlet and the Grado inlet (Figs. 1 and 4, Table 2).

Singlebeam sonar data was acquired along the coast of the Tagliamento delta between the Baseleghe and Lignano inlets, from the Morgo inlet to the Marina Julia beach and in a dense grid in front of the Porto Buso inlet (Figs. 1 and 4, Table 2).

Single channel seismic and acoustic data were acquired in various surveys (Table 2 and Fig. 4). Interpretation of the geophysical profiles was done by means of the IHS Kingdom® software. We used seismic and acoustic profiles to determine the seafloor depth where singlebeam or multibeam bathymetric data were not acquired. The seismic and acoustic profiles used to create the bathymetric model stretch over a large part of the western part of the gulf, offshore Grado, southeast of Miramare and in the Bay of Muggia (Figs. 1 and 4).

The multibeam and singlebeam sonar data coupled with the interpreted depths of the seafloor from acoustic and seismic profiles were imported in the SKUA-GOCAD™ Paradigm software package in order to perform quality control and consequent adjustments to the dataset. In order to convert the seismic and acoustic data from the time to the depth domain, 1,514 m/s was assumed as the velocity of sound in the water column. This value



represented the average value of sound velocity profiler measurements acquired together with the datasets used in this study. After the conversion gridding of the bathymetric model was done with the ‘Discrete Smooth Interpolation’ (DSI) method (Mallet, 1992, 1997) in the SKUA-GOCAD™ software package. The coordinates of the pointsets building the model were clipped to the coastline extent and later exported as a column based text file and a geotiff grid with a 50 m x 50 m cell size. The resulting model contained some areas with negative depth values that resulted from lack of geophysical data in the near proximity of the coast. In addition, the morphologically very dynamic northern part of the research area starting from the Isonzo mouth towards the west also contained negative depth values due to transient sedimentary bodies elevated above or just below the sea-level that were present at the time of the various surveys but were not delineated by the coastline shapefile used to clip the dataset. In order to still represent these dynamic shallow areas while eliminating artificial negative depths from our dataset, we replaced all the negative values in the geotiff grid and the column based text file with a depth value of 0.2 m.

## 2.2 Model of the base of the Holocene

Seismic (boomer) and acoustic data (Chirp and parametric sub-bottom sonar) were used in order to create a model of the base of the Holocene marine sediment in the Gulf of Trieste (Fig. 5 and Table 3). In general, the used profiles span throughout the Gulf of Trieste but are also located in the inner parts of the Grado Lagoon, Bay of Muggia and in the Bay of Koper (Fig. 5). The base of the Holocene was determined from reflection characteristics visible from profiles that were interpreted by means of the IHS Kingdom® software and calibrated with the cores listed in Table 1. In general, the base of the Holocene marine sediment on geophysical profiles can be recognised as a prominent reflector underlying the acoustically transparent marine sediment (Figs. 3 and 8; Slavec, 2011, 2012; Trobec, 2015; Trobec et al., 2017). Additionally, depths of the base of Holocene marine sediment determined from offshore and onshore well and core data from Italy and Slovenia were included in the model and are shown in Table 1 and Figs. 2 and 5.

This dataset was imported, quality controlled and adjusted in the SKUA-GOCAD™ software package. A velocity of 1,530 m/s was assumed for the velocity of sound traveling through the Holocene sediment and used for the conversion of seismic and acoustic data from time to depth domain. Afterwards, gridding of the model of the base of the Holocene was done in the SKUA-GOCAD™ software package with the DSI interpolation method. Even though the sound velocity of the sediment overlying the flysch in the Gulf of Trieste was determined at 1,610 m/s (Masoli et al., 2015), we decided for a lower value because Masoli et al. (2015) considered a few ten meters thick Late Pleistocene – Holocene sedimentary sequence where compaction most probably already affected the velocity value. Considering that: a) the typical sound velocity of a marine water saturated sediment is approximately 1,500 m/s (Anderson and Hampton, 1980; Yuan et al., 1992); b) the average value of sound velocity profiler measurements acquired together with the datasets used in this study is 1,514 m/s; c) that the Holocene marine sediment thickness in our study is an order of magnitude smaller than the sediment considered in Masoli et al. (2015), leads us to believe that our chosen velocity is reasonable for our conversion. The coordinates of the pointsets building the model were clipped to the coastline extent and later exported as a column based text file and a geotiff grid with a 50 m x 50 m cell size.



## 2.3 Modeling the thickness of Holocene marine sediment

5 A model of the thickness of Holocene marine sediment in the Gulf of Trieste was created with the SKUA-GOCAD™ software. In order to create the model, the difference between the bathymetric model and the model of the base of the Holocene was calculated. Data used to create the model of the base of the Holocene spanned over a smaller area compared to the bathymetric model (Figs. 4, 5, 6 and 7). In areas where data coverage was insufficient or even absent (grey areas bounded by black dashed lines on Fig. 5), we used data from the ‘Carta degli spessori dei sistemi di stazionamento alto (HST)’ published by Trincardi et al. (2011b), for which we  
10 assumed 1,530 m/s as the velocity of sound traveling through the sediment. We excluded the areas inside the Grado-Marano lagoon from the model due to dubious interpretation of geophysical profiles (signal reverberation) resulting from shallow seafloor depth and because these Holocene sediments were deposited in a lagoonal, rather than a marine sedimentary environment. The coordinates of the pointsets building the model  
15 were clipped to the coastline extent and later exported as a column based text file and a geotiff grid with a 50 m x 50 m cell size. The model contained negative thickness values in areas of erosion (e.g. Fig. 8(e)) and in the western part of the research area, where Holocene marine sediment is very thin or even absent and is therefore very difficult to model. In this area thicker sedimentary sequences can seldom be found only as infill of Late Pleistocene channels (Figs. 8(a) and (c)). When modelling such geometries, the base of the Holocene surface  
20 tends to rise above the seafloor outside the channel as a convex more or less pronounced bulge. This in turn results as artificial negative thickness values. In order to honour the data, we replaced all the negative values in the geotiff grid and the column based text file with a thickness value of zero.

## 25 3 Results

### 3.1 Bathymetry of the Gulf of Trieste

30 The bathymetry of the seabed of the Gulf of Trieste has a generally smooth morphology. Depth values in our model vary between 0.2 and 32.6 meters below sea level with a mean of 16.2 meters and standard deviation of 6.2 (Fig. 6). In the southeastern and eastern parts of the gulf depth exceeds the 20 meter isobath approximately 4 kilometres away from the shore at most, while in the northern and northwestern parts the distance needed is approximately twice as much (but can reach up to 20 kilometres).

35 The Trezza Grande paleodelta (Gordini et al., 2002; Zecchin et al., 2015) and the dune shaped features NW of Piran (Slavec, 2012) are topographically higher areas visible on Fig. 6. The deepest feature on the bathymetric model is the elongate Cape Madonna depression stretching in a SW-NE direction north of Piran (Fig. 6) which also represents the deepest point of the Gulf of Trieste with a depth of 38 meters (Ministry of Transport of the Republic of Slovenia, 2005; Slavec, 2012). A deeper area where depths exceed 25 meters is located north of the  
40 Piran and west of the Bay of Muggia.



### 3.2 Base of the Holocene marine sediment

- 5 Contrary to the smoothness of the bathymetric model, the model of the base of the Holocene marine sediment presents rougher morphologies (Fig. 7), which are evident also from Figs. 3 and 8. The model of the base of the Holocene is located between 6.6 and 31.0 meters below sea level with a mean of 19.8 meters and a standard deviation of 5.7. In general, the model can be divided in two parts, the shallower northwestern and the deeper southeastern part (Fig. 7). This difference in depth is also evident from geophysical profiles (Figs. 8(a) and (b)) and core data (Fig. 2).
- 10 The northwestern part is shallower than 20 meters and in general rises as we approach the coastline. Local topographical highs are located in the southern part of the Trezza Grande area (Fig. 7). In places the model exhibits channel-like features that are a few meters deeper from the surrounding topography, but their exact course is difficult to map due to insufficient profile coverage.
- 15 The base of Holocene marine sediment in the southeastern part of the Gulf forms a basin that is in general deeper than 25 meters below sea level (Fig. 7) and is also visible on geophysical profiles and recognisable in core data (Figs. 8 and 2). Shallower depths can only be observed in the Bay of Muggia and in the immediate vicinity of the coastline where it rises rapidly on a very short distance (Fig. 7). A slightly deeper section oriented in a WNW-ESE direction approximately 6 km north of Piran is evident and corresponds to the meander belt of the Paleorizana (Slavec, 2012; Trobec et al., 2017).
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### 3.3 Thickness of Holocene marine sediment

- 25 The modelled thickness of the Holocene marine sediment in the Gulf of Trieste extends between 0 and 24.0 meters with a mean thickness of 3.2 meters and a standard deviation of 2.8. In general, the inner parts of the gulf are covered by very thin drapes of Holocene marine sediment that can also be partially absent (Figs. 9, 8(a) and (b)). Towards the shore, the Holocene sedimentary sequence gradually thickens and appears as a coastal sedimentary wedge that can exceed thicknesses of ten meters near the coastline and in the internal parts of bays.
- 30 The map of the thickness of the Holocene marine sediment in the Gulf of Trieste can be divided into two parts, the central – western part with thinner marine sequences and the southeastern part of the gulf with thicker Holocene marine sediments (Fig. 9). This prominent difference is clearly evident from geophysical profiles spanning through a larger part of the Gulf (Figs. 8(a) and (b)) and from a simplified W-E stratigraphic profile from the published cores (Fig. 2).
- 35 The central - western part of the gulf is characterized by a very thin Holocene sedimentary cover which rarely exceeds a thickness of three meters (Figs. 9, 8(a) and (b)). Where the thin sediments are not well resolved or even unresolvable with the high resolution geophysical methods used in this study, we assume they are absent or only a few centimetres thick. Thicker Holocene sedimentary sequences can be observed only in the vicinity of the Tagliamento delta (Figs. 9, 3 and 4 from Zecchin et al., 2015), the Trezza Grande paleodelta (Figs. 2 and 9;



Figs. 4 and 5 from Zecchin et al., 2015) and in the central part of the bay where thicker sedimentary sequences fill pre-Holocene paleochannels (Figs. 9, 7, 8(a), (b) and (c)).

The southeastern part of the gulf is characterised by a thicker sedimentary cover which averages at about 5 meters (Figs. 9, 2, 8(a) and (b)). Further examination reveals that in some places in close proximity of the shoreline the Holocene sedimentary cover often thins out or is unresolvable on geophysical profiles (for example near Debeli rtič, Fig. 8(d)). This effect is a consequence of basement rock (predominantly flysch) outcropping and/or subcropping near the seafloor. In the Cape Madonna depression in front of Piran (Fig. 6) the marine Holocene sedimentary sequence is not recognisable due to erosion which can also be identified on geophysical profiles down to the Late Pleistocene sediment (Fig. 8(e)). In addition to thick marine Holocene sedimentary deposits in the internal parts of bays (Figs. 9 and 2), thicker sedimentary sequences can be observed over buried paleochannels (Figs. 9, 8(a), (b) and (c); Slavec, 2012; Trobec, 2015; Trobec et al., 2017) and in the area of the dune shaped features NW of Piran (Figs. 6, 8(f) and 9).

#### 4 Model quality assessment

While our work was done with great care and attention to detail in order to produce accurate results, we would like to emphasize that the model is primarily intended to illustrate the general trend of the thickness of Holocene marine sediment in the Gulf of Trieste and cannot account for strong local thickness variations. This is a result of several factors that are described in the following section.

As mentioned in Sect. 3.2 and 3.3, the surface of the base of the Holocene marine sediment can be very undulated due to the pre-existing paleotopography and/or due to subcropping or outcropping of the basement rock on the seafloor near the eastern and southern coastline (Fig. 8(d)). Combined with areas where density of acoustic and seismic profiles is not adequate, this undulation can lead to artefacts that are later outstretched throughout the model of the base of the Holocene and consequentially affect the values of the model of the thickness. Discrepancies of the map with the natural state are also a result of the gridding algorithm used (Mallet, 1992, 1997). Due to the inherent two-way travel time attribute of the acquired data in the case of geophysical profiles, a major influencing factor on the models are the chosen velocity for the conversion to the time domain and the local sound velocity variations in the Holocene sediment which are unknown. As there are no measurements of speed of sound traveling solely through the Holocene sediment of the Gulf of Trieste, we used a conservative approximation of the speed of sound traveling through the sedimentary column, which could lead to systematic underestimation of the thickness throughout the whole study area. Due to a lack of sound velocity measurements we also assumed a model without velocity variations. Finally, due to scarce cores/wells, the presence of marine sediment in the Holocene sedimentary sequence was assessed mostly by following visual criteria when investigating geophysical profiles. We considered the acoustically transparent Holocene seismic facies to represent sediment primarily deposited in a marine environment. In order to avoid including predominantly paralic Holocene sediment in the model, we excluded the Grado-Marano lagoon area from the model of the thickness of Holocene marine sediment.



## 5 Data availability

An important moment of data lifecycle is that when observations are shared with other team members or domain experts in general. This introduces the possibility to replicate the experiment, to re-use the data in other contexts, and to provide data with additional value. OGS is deeply involved in the curation and dissemination of its data assets and has in this perspective developed a web based data system called SNAP (Diviacco et al 2015, Diviacco and Busato 2013) which grants that the contained data are findable, accessible, interoperable and reusable (FAIR). The system is based on fully compliant OGC O&M and SensorML metadata and offers previewing facilities that allow remote data access and download once the end user checked whether they are what he/she is looking for. DOIs are assigned to the datasets residing in the system so that they can be directly accessed resolving the DOIs or simply following the corresponding URLs. In the case of the models used in this study, they are accessible at the following URL: <https://dx.doi.org/10.6092/6ad9b1e6-c977-cec9-8a2d-db10c7f90adc>. This points to a landing page which has been built following the standards set in the UNESCO Ocean data Publication Cookbook (Leadbetter et al., 2013), and offers further metadata details and an interactive preview of the dataset. To download data, a registration to the system is required.

The data described in this study are available as column based text files containing data points with X, Y and Z/thickness coordinates in the UTM 33 North coordinate system (datum: WGS 84). In order to facilitate their use (especially with GIS software) these datasets are also available as 50 x 50 m elevation grids clipped to the extent of the coastline in the widely compatible Geotiff format along with georeferencing information (the \*.proj and \*.tfw files). Depth contours of the models are also available in the shapefile format.

## 6 Conclusions

Together with the published datasets (<https://dx.doi.org/10.6092/6ad9b1e6-c977-cec9-8a2d-db10c7f90adc>), this paper represents the second assessment of the thickness of Holocene marine sediment in a major part of the Gulf of Trieste with employing new datasets from a wide range of different geophysical methods and published sedimentological data from wells and cores.

Geophysical profiles along with core/well data show a basin in the Late Pleistocene topography in the southeastern part of the Gulf that served as a depocentre for the significantly thicker Holocene marine sedimentary sequences compared to the rest of the gulf. While the exact cause for the basin formation exceeds the scope of this paper, we suspect it represents an interplay between different tectonic, sedimentologic, climatic and oceanographic factors that need to be further examined in order to better understand the distribution of Holocene marine sediment in the Gulf of Trieste. Additionally, thicker Holocene marine sequences can be observed above paleochannels, in the Trezza Grande area and in the internal parts of bays and in close proximity of the shoreline, where the thickness is probably governed by increased amounts of available sediment transported by rivers (Covelli et al., 2004; Mandac Soczka and Faganeli, 2015) and surface run-off from rocky coasts of the southeastern part and due to increased amounts of fluvially transported sediment from the alluvial plains of the northern and northwestern coasts of the gulf. In the central part of the Gulf the Holocene marine sequences are very thin or even absent (unresolvable on geophysical profiles).





Our work provides a solid reference for a wide variety of disciplines involved in future studies of Holocene sediment in the area, especially regarding sampling sites and/or survey selection. Furthermore, the thickness model provides important implications for further sedimentological, geomorphological, paleoenvironmental, neotectonic and oceanographic studies of the Gulf of Trieste.

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### Author contributions

AT interpreted acoustic data from the Slovenian part of the gulf, modeled the surfaces and thickness, prepared the figures and wrote the manuscript. MB conceived and managed the study, interpreted the geophysical data from the Italian part of the Gulf and contributed in writing the manuscript. FZ, LB, AC, EG, RR and IT acquired the multibeam, singlebeam and sonar data in various campaigns in the Italian part of the Gulf. AB, EG and RR interpreted various geophysical datasets in the Italian part. LB and FZ processed the geophysical data from the Italian part. SP acquired and processed the multibeam dataset and the sub-bottom sonar profiles from the Slovenian part of the gulf. PD managed data and metadata handling on the online repository. MV interpreted acoustic data from the Slovenian part of the gulf.

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### Competing interests

The authors declare that they have no conflict of interest.

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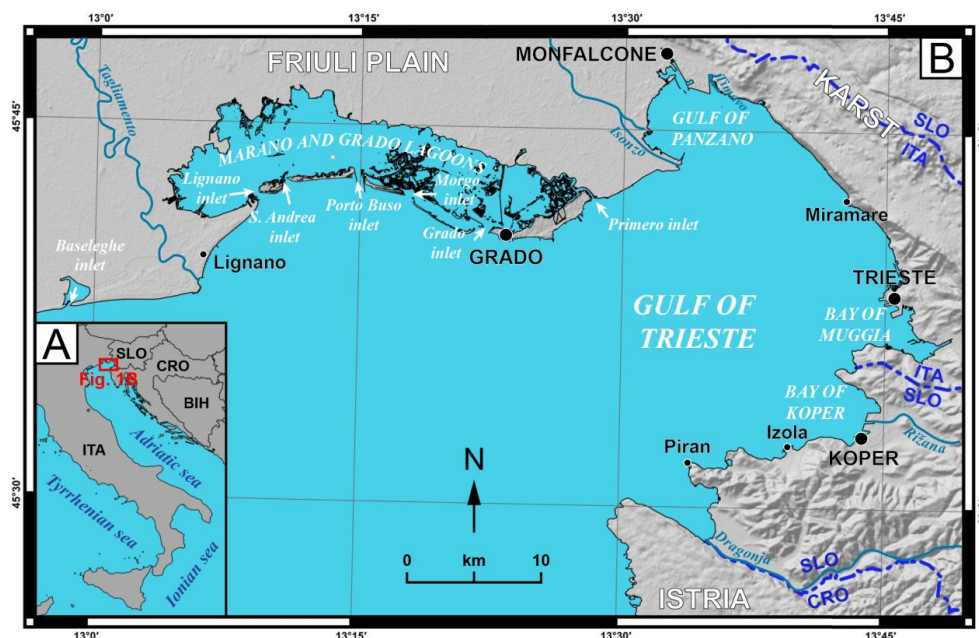
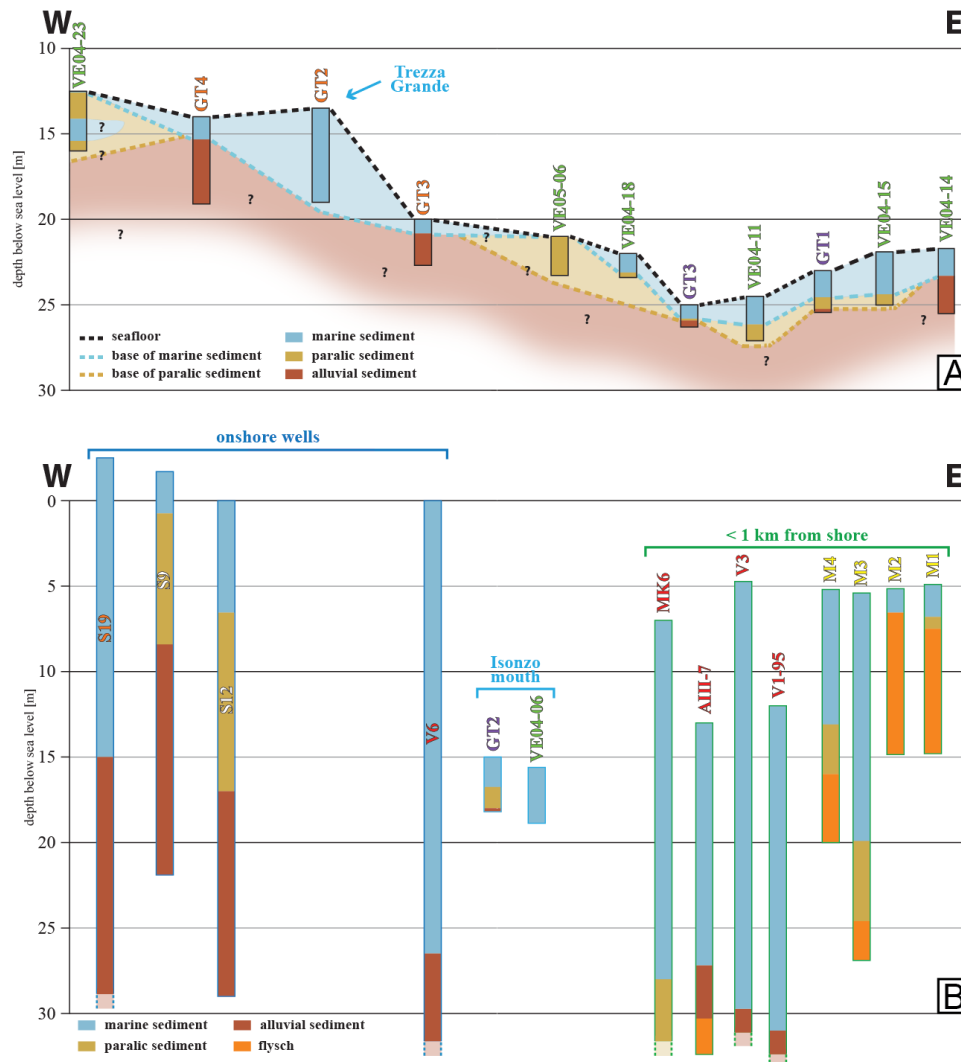


Figure 1: Location of the study area; (A) Regional map; (B) Geographic surroundings of the study area.



**Figure 2: Simplified stratigraphy of all the cores and wells used in this study** (Table 1; Ogorelec et al., 1981, 1991, 1997; Marocco et al., 1984; Marocco, 1989; Gordini et al., 2002; Covelli et al., 2006; Romeo, 2009; Trincardi et al., 2011b; Zecchin et al., 2015). Note that the distance between the cores/wells is not in scale; (a) A simplified W-E stratigraphic profile of the cores located in the central part of the Gulf of Trieste; (b) Wells and cores used in the study that were not suitable for the stratigraphic profile due to their location or proximity to the shore or fluvial sedimentary sources (see Figs. 4 and 5 for location). Wells outlined in dark blue are located onshore. Cores outlined in bright blue are located in close proximity of the present-day Isonzo river mouth. Cores outlined in green are located less than 1 kilometre from the shoreline. Dashed bottoms of wells indicate that their whole stratigraphy is not illustrated in the figure and that the reader should refer to the original publications for further information.



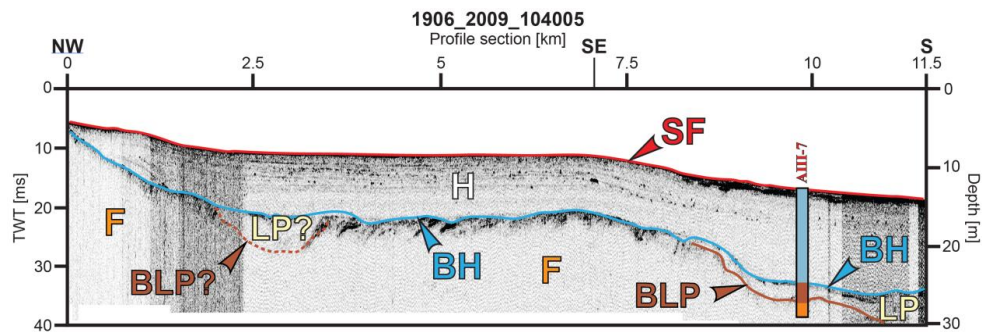


Figure 3: An example of an acoustic profile correlated with the core AIII-7 from Ogorelec et al., 1997 (SF – seafloor, H – Holocene marine sediment, BH – reflector that marks the base of the Holocene sediment, LP – Late Pleistocene sediment, BLP – base of the Late Pleistocene sediment, F – flysch, TWT – two-way travel time).

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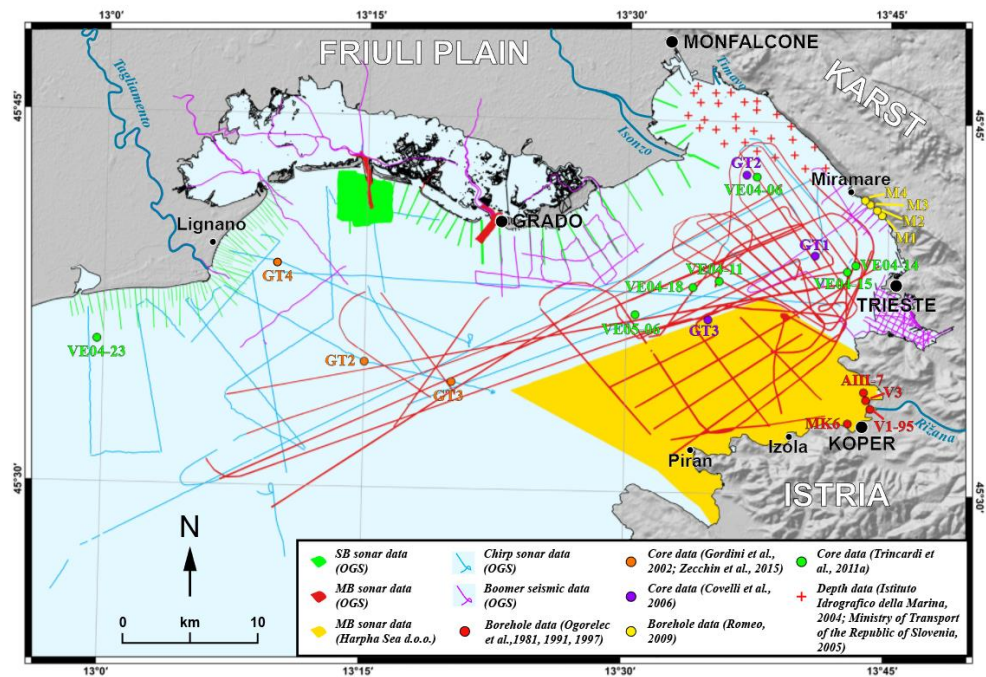


Figure 4: Locations of the datasets used for the bathymetric model of the Gulf of Trieste.



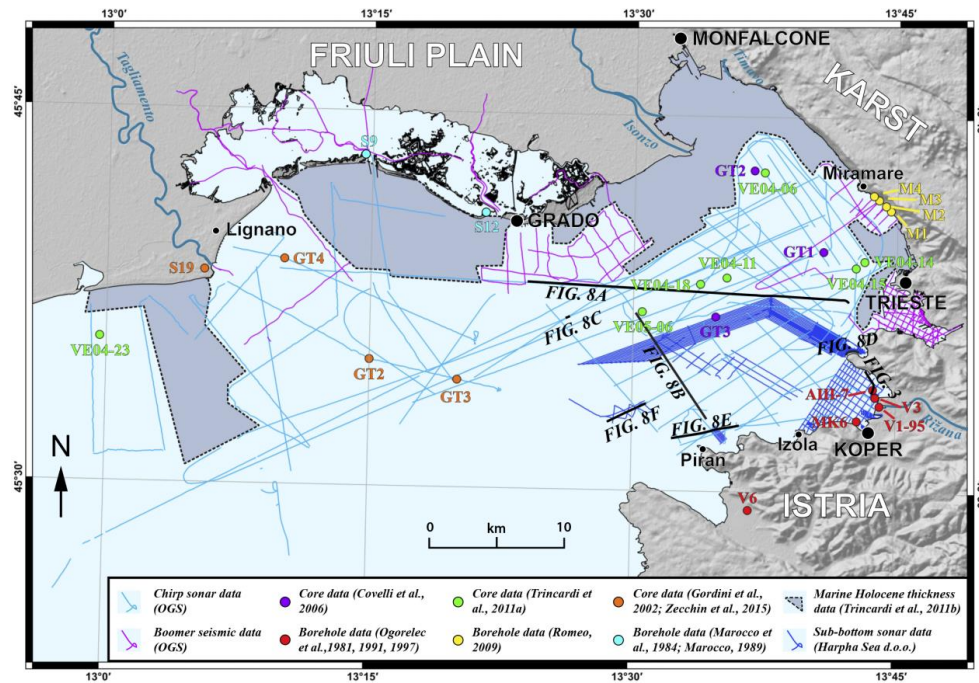


Figure 5: Locations of the datasets used for the model of the base of the Holocene. Black lines indicate locations of acoustic profiles shown in Figs. 3 and 8. Note that the black dashed lines mark areas where published data from Trincardi et al. (2011b) was used for the model of the thickness of Holocene marine sediment.

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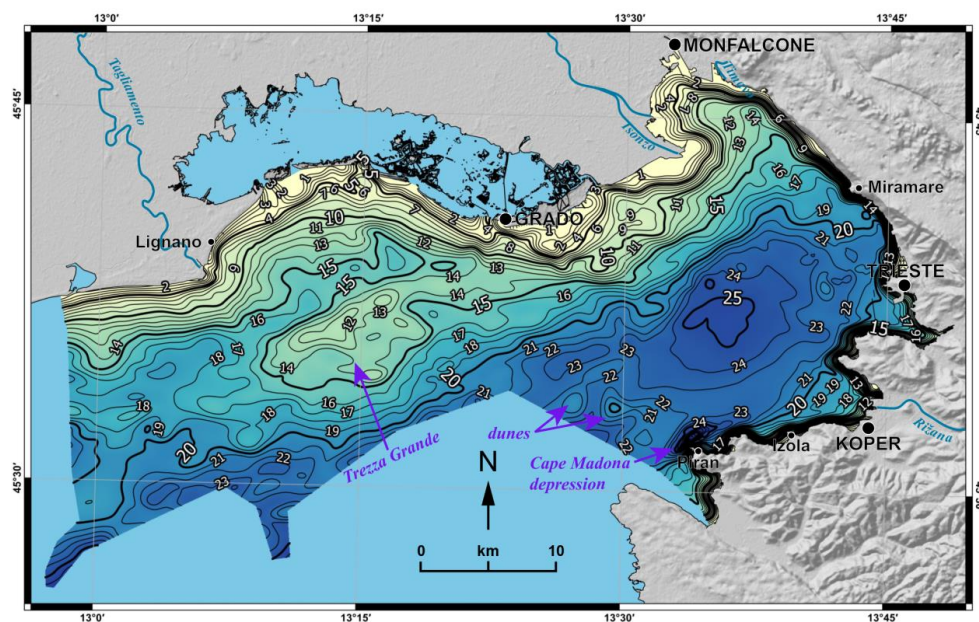


Figure 6: Map of the bathymetry of the Gulf of Trieste with indicated prominent morphological features.

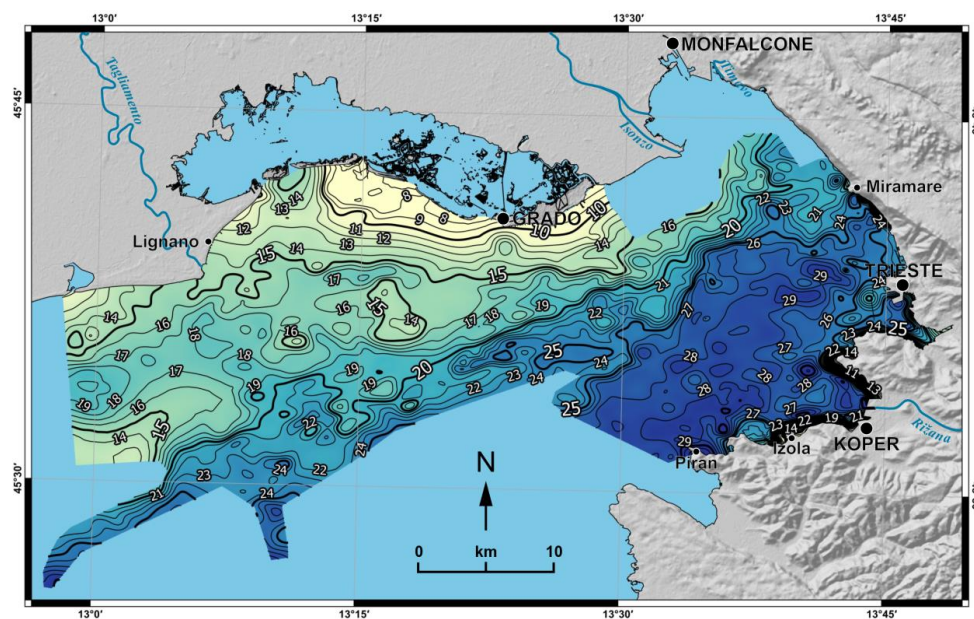


Figure 7: Map of the base of the Holocene marine sediment.

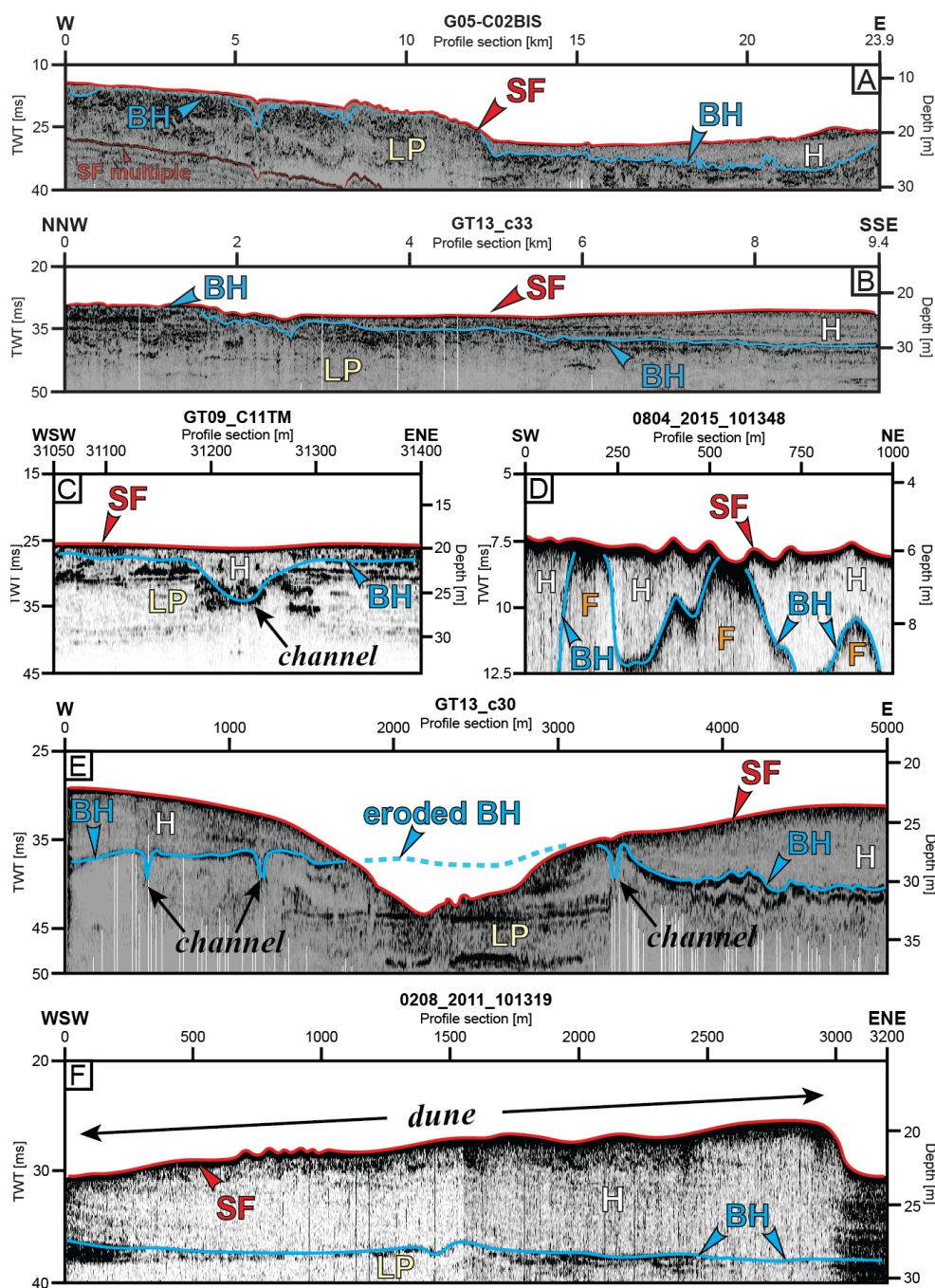


Figure 8: Examples of sub-bottom and Chirp sonar profiles (for explanation of symbols refer to caption of Fig. 3); (a-b) Regional Chirp profiles show that the Late Pleistocene paleotopography forms a basin in the eastern part which serves as a depocentre for the Holocene marine sediment (dashed blue line – base of the Holocene sediment where the sedimentary sequence is very thin or even absent or eroded); (a) A W-E oriented profile showing a difference in

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marine Holocene thickness in the western and eastern part of the Gulf. Note that thicker Holocene marine sequences in the western part of the Gulf are filling channel-like features; (b) A NNW-SSE oriented profile showing the same difference in Holocene marine thickness; (c-f) Variations of the thickness of marine Holocene sediment; (c) A profile from the central part of the Gulf, where a thin Holocene marine sediment cover is characteristic, except where it fills pre-existing channels; (d) Variation of Holocene sediment thickness controlled by the pre-existing Flysch paleotopography in the area of Debeli rtič; (e) Holocene sediment erosion in the seafloor depression in front of Piran (Fig. 6), note also Holocene filling of pre-existing channels; (f) Thicker Holocene sedimentary sequences visible on a profile crossing the area of dunes in front of Piran (Fig. 6).

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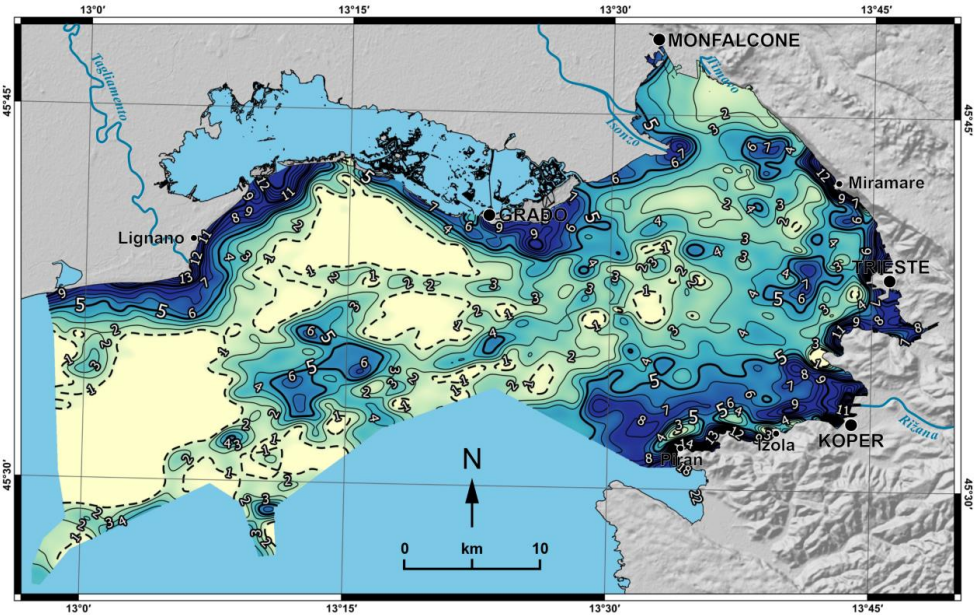


Figure 9: Figure 7: Map of the thickness of the Holocene marine sediment. Dashed black lines mark areas where Holocene marine sediment is thinner than one meter or even absent.

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Core/well	Water depth [m]	Holocene thickness [m]	Reference	Label colour
VE04-06	15.6	>3.27	Trincardi et al., 2011a	green
VE04-11	24.5	1.63	Trincardi et al., 2011a	green
VE04-14	21.7	1.61	Trincardi et al., 2011a	green
VE04-15	21.9	2.48	Trincardi et al., 2011a	green
VE04-18	22	1.12	Trincardi et al., 2011a	green
VE04-23	12.5	0.05	Trincardi et al., 2011a	green



VE05-06	21	0	Trincardi et al., 2011a	green
S12	0	6.55	Marocco et al., 1984	blue
S9	-1.7	2.45	Marocco, 1989	blue
GT2	13.5	>5.43	Gordini et al., 2002 & Zecchin et al., 2015	orange
GT3	20	0.83	Gordini et al., 2002 & Zecchin et al., 2015	orange
GT4	14	1.32	Gordini et al., 2002 & Zecchin et al., 2015	orange
S19	-2.5	17.5	Gordini et al., 2002 & Zecchin et al., 2015	orange
GT1	23	1.55	Covelli et al., 2006	purple
GT2	15	1.75	Covelli et al., 2006	purple
GT3	25	0.85	Covelli et al., 2006	purple
M1	4.9	1.9	Romeo, 2009	yellow
M2	5.15	1.4	Romeo, 2009	yellow
M3	5.4	14.5	Romeo, 2009	yellow
M4	5.2	7.9	Romeo, 2009	yellow
AIII-7	13	14.2	Ogorelec et al., 1997	red
V1-95	12	19	Ogorelec et al., 1997	red
MK6	7	21	Ogorelec et al., 1991, 1997	red
V3	4.5	25	Ogorelec et al., 1991, 1997	red
V6	0	26.5	Ogorelec et al., 1981	red

**Table 1: Wells and cores used in this study. The > symbol indicates, that the core contains only Holocene marine sediment.**

Area	Year	Institute	Type	Instrument	Quantity	Reference
Porto Buso inlet	2006	OGS	Multibeam	Reson Seabat 8125	1.6 km <sup>2</sup>	Cova, 2008
Grado inlet	2007	OGS	Multibeam	SWATHplus-Wide swath Bathymetry & Side scan sonar system Sea	1.7 km <sup>2</sup>	Gordini, 2008a
Morgo inlet	2007	OGS	Multibeam	SWATHplus-Wide swath Bathymetry & Side scan sonar system Sea	0.3 km <sup>2</sup>	Gordini, 2008a
Central part of the Gulf	2009	OGS	Multibeam	Reson SeaBat 8111 hull mounted	49 km <sup>2</sup>	Zgur et al., 2010
Slovenian and Central part of the Gulf	2013	OGS	Multibeam	Reson SeaBat 8111 hull mounted	20.6 km <sup>2</sup>	Zgur et al., 2013
Slovenian part of the Gulf	2006-2009	Harpha Sea	Multibeam	Reson SeaBat 8125 & Elac HydroStar	205 km <sup>2</sup>	Poglajen, 2012; Slavec,



		d.o.o.		4300		2012
Porto Buso – Monfalcone	2007	OGS	Singlebeam	Garmin echosounder	62 km	Gordini, 2008b
Porto Buso	2007	OGS	Singlebeam	Garmin echosounder	250 km	Gordini, 2008c
Tagliamento Delta	2002-2003	OGS	Singlebeam	Lowrance LCX – 18C echosounder	120 km	Gordini et al., 2006

Table 2 – Multibeam and singlebeam datasets used for the bathymetric model.

Area	Year	Institute	Type	Instrument	Quantity	References
Lignano	2000	OGS	Boomer	Boomer CEA PULSAR 2002 & UWAK 05	61.7 km	Della Vedova et al., 2002
Lignano	2002-2003	OGS	Chirp	Datasonic Sub-bottom Profiler CHIRP CAP- 6600	153.1 km	Gordini, 2009; Zecchin et al., 2015
Grado & Grado- Marano lagoon	2008	OGS	Boomer	Boomer CEA PULSAR 2002 & UWAK 05	60.9 km	Baradello et al, 2009
Miramare	2003	OGS	Boomer	Boomer CEA PULSAR 2002 & UWAK 05	16.2 km	Romeo, 2009
Gulf of Trieste	2005	OGS	Chirp	Benthos CHIRP II Hull mounted 16 transducers	179.2	Busetti et al., 2005
Gulf of Trieste	2009	OGS	Chirp	Benthos CHIRP II Hull mounted 16 transducers	516.4 km	Zgur et al., 2010
Gulf of Trieste	2013	OGS	Chirp	Benthos CHIRP II Hull mounted 16 transducers	160.3 km	Zgur et al., 2013
Muggia	2013	OGS	Boomer	Boomer CEA PULSAR 2002 & UWAK 05	76.5 km	Baradello et al., 2013
Piran-	2013	OGS	Boomer	Boomer CEA	26.1 km	Baradello,



Koper				PULSAR 2002 & UWAK 05		pers. comm.		
Slovenian waters	2009-2015	Harpha d.o.o.	Sea	Parametric sub-bottom sonar	Innomar Sediment Sounder Compact	Parametric Echo SES-2000	560.8 km	Žerjal et al., 2015, Poglajen, pers. comm.
Strunjan bay	2013-2014	Harpha d.o.o.	Sea	Parametric sub-bottom sonar	Innomar Sediment Sounder Compact	Parametric Echo SES-2000	53.3 km	Trobec, 2015

**Table 3: Seismic and acoustic data used for the model of the base of the Holocene marine sediment.**